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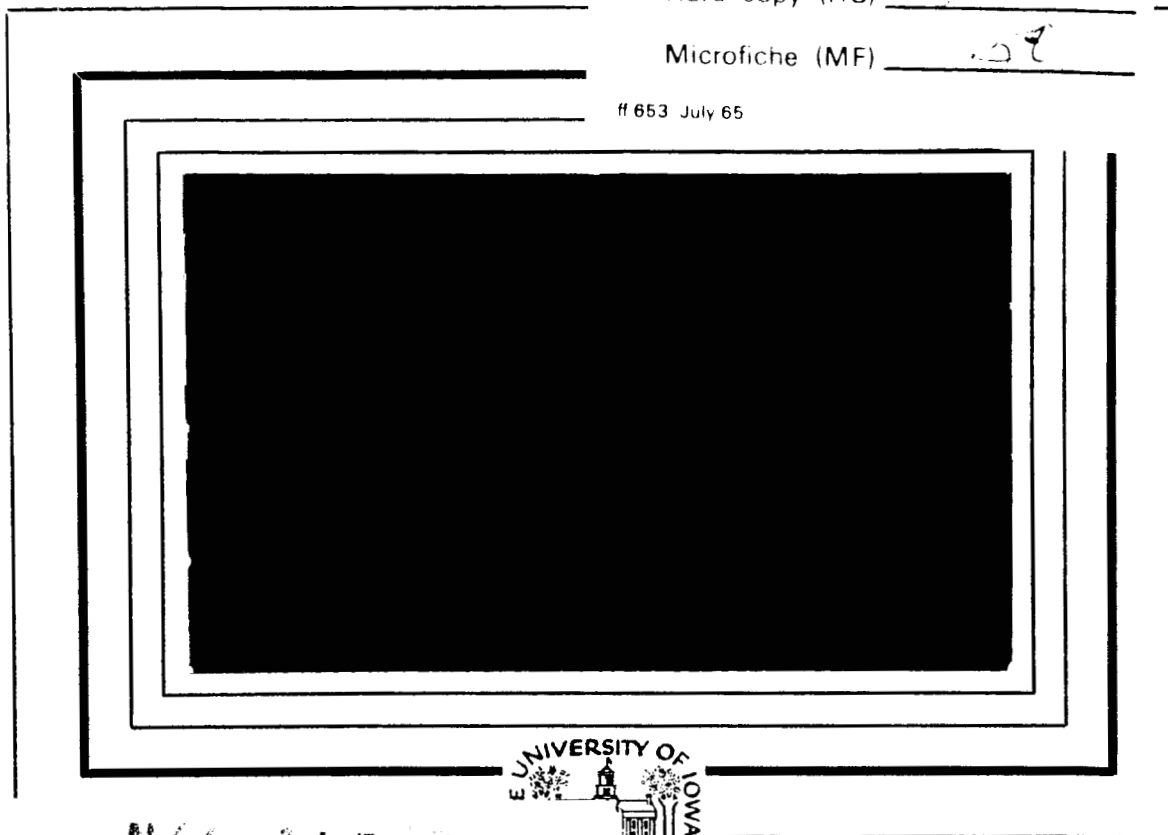
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Determination of Hydrogen Ion  
Concentration, Electron Density  
and Proton Gyrofrequency from  
the Dispersion of Proton Whistlers \*

by

Donald A. Gurnett  
Stanley D. Shawhan \*\*

Department of Physics and Astronomy  
University of Iowa  
Iowa City, Iowa

\* Research supported in part of Office of Naval Research  
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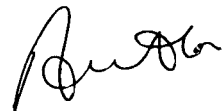
\*\* National Aeronautics and Space Administration Graduate  
Trainee

## ABSTRACT

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In this paper we discuss a method for determining  $H^+$  concentration, electron number density, and proton gyrofrequency in the vicinity of the satellite by measurements of the asymptotic frequency-time profile of a proton whistler near the proton gyrofrequency. This new technique is applied to proton whistlers received by the Injun 3 VLF receiver. The calculated values of  $H^+$  concentration and electron density are shown to be in good agreement with measurements by other experimenters at similar altitudes, latitudes, and local times. B values calculated from the proton gyrofrequency are compared with values calculated from the Jensen and Cain expansion for the geomagnetic field.

It is shown that the wave energy of a proton whistler is guided very nearly along the geomagnetic field and that the parallel component of the group velocity is closely approximated by the group velocity for longitudinal propagation. It is found that for frequencies near the proton gyrofrequency at the satellite the kernel multiplying  $n(H^+)^{1/2}$  in the travel time integral is large only in the region near the satellite. Assuming that the  $H^+$  concentration is uniform within this region, an expression is derived for the travel time of a proton whistler near the proton gyrofrequency.



The  $H^+$  concentration and proton gyrofrequency are obtained by fitting this theoretical frequency-time expression to observed proton-whistler signals. By combining this method for determining  $n(H^+)$  with the crossover frequency method for determining  $\alpha_1 = n(H^+)/n_e$  the electron density can also be determined.

## I. INTRODUCTION

In the VLF recordings from the Injun 3 and Alouette 1 satellites, an unusual ion gyrofrequency phenomenon was observed following the reception of short fractional-hop whistlers [Smith et al., 1964]. The new effect appears as a tone which starts immediately following the reception of a short fractional-hop whistler and initially shows a rapid rise in frequency, asymptotically approaching the gyrofrequency for protons in the plasma surrounding the satellite (see Figure 1). The phenomenon was explained [Gurnett et al., 1965 and Gurnett, 1965] as simply a dispersed form of the original lightning impulse with the dispersion arising from the effects of ions on the propagation. The characteristic tone occurring immediately following the reception of a whistler was identified as a left-hand polarized ion cyclotron wave and is called a proton whistler. A detailed discussion on the occurrence of proton whistlers in the Injun 3 data is given by Shawhan [1965].

At a frequency denoted by  $\omega_{12}$ , the proton whistler and electron whistler frequency-time traces are coincident in time (see Figure 1). This frequency was identified as the crossover frequency, a characteristic frequency of the plasma surrounding the satellite. The abrupt change in the group velocity near  $\omega_{12}$  occurs as the wave polarization changes sign at the crossover frequency. Measurements of the crossover frequency using proton whistlers received

by Injun 3 have been used to provide accurate measurements of the fractional concentration  $\alpha_1 = n(H^+)/n_e$  in the ionosphere [Shawhan and Gurnett, 1965].

In this paper we discuss a method of determining the  $H^+$  concentration, electron number density, and proton gyrofrequency in the vicinity of the satellite by measurements of the travel time of a proton whistler for frequencies near the proton gyrofrequency. This technique is applied to proton whistlers received by Injun 3. The calculated values of  $H^+$  concentration and electron density are shown to be in good agreement with measurements by other experimenters at similar altitudes and local times. Values of the geomagnetic field computed using this proton whistler method are compared with values obtained from the Jensen and Cain spherical harmonic expansion of the geomagnetic field for epoch 1960.

This new method of determining the local  $H^+$  concentration, electron density, and proton gyrofrequency is based on the theoretical explanation of the proton whistler. As discussed by Gurnett et al. [1965] and Gurnett [1965], the proton whistler is a left-hand polarized ion cyclotron wave propagating in the frequency band available just below the proton gyrofrequency. For frequencies near an ion gyrofrequency the dispersion relation for the ion cyclotron wave becomes particularly simple and is discussed in Section 2 of this paper. It is shown that the wave energy of a

proton whistler is guided very nearly along the geomagnetic field and that the group velocity is nearly independent of the wave normal angle and can be closely approximated by the group velocity for longitudinal propagation. The group travel time of proton whistlers observed at the satellite is given by a line integral along the ray path from the lightning source to the satellite. It is found that for frequencies near the proton gyrofrequency (at the satellite) the kernel multiplying  $n(H^+)^{1/2}$  in the integrand is large only in the region near the satellite. Assuming that the  $H^+$  concentration is approximately uniform within this region, an expression is derived for the travel time of the proton whistler for frequencies near the proton gyrofrequency. The  $H^+$  concentration and proton gyrofrequency in the vicinity of the satellite are obtained by fitting this theoretical travel time expression to observed frequency-time traces for proton whistlers. By combining this method for determining  $n(H^+)$  with the crossover frequency method for determining  $\alpha_1$  the electron density  $n_e$  can be calculated using  $n_e = n(H^+)/\alpha_1$ .

## II. THE GROUP TRAVEL TIME NEAR AN ION GYROFREQUENCY

The group travel time of a proton whistler from the source of the lightning impulse to the satellite is given by the line integral:

$$t(\omega) = \int_p \frac{ds}{u} \quad (1)$$

$$u = \left| \frac{\partial \omega}{\partial \vec{k}} \right|$$

Before this integral can be evaluated, we must consider the ray path  $p$  and the group velocity  $u$  along the ray path. Since the group velocity of the left-hand polarized ion cyclotron wave is much less than the group velocity of the right-hand polarized wave, we recognize that the major contribution to the travel time integral occurs along the portion of the path where the wave is left-hand polarized. This is especially true for frequencies near an ion gyrofrequency since the group velocity of the ion cyclotron wave approaches zero at the ion gyrofrequency. Thus, we limit our attention to an ion cyclotron wave with frequencies near an ion gyrofrequency.

Using the notation introduced by Stix [1962] the refractive index for ion cyclotron waves can be written\*:

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\*Subscripts (k) are e, 1, 2, and 3 for electrons,  $H^+$ ,  $He^+$ , and  $O^+$  respectively



$$n^4 \cos^2 \theta - n^2 S(1 + \cos^2 \theta) + RL = 0 \quad (2)$$

$$R = 1 - \sum_k \frac{\Pi_k^2}{\omega(\omega + \Omega_k)} \quad (3)$$

$$L = 1 - \sum_k \frac{\Pi_k^2}{\omega(\omega - \Omega_k)} \quad (4)$$

$$S = (1/2)(R + L)$$

The plasma frequency and gyrofrequency of the  $k$ th constituent are  $\Pi_k$  and  $\Omega_k$  (including sign) respectively, and  $\theta$  is the angle between the wave normal and the static magnetic field. In the derivation of equation (2), it is assumed that  $L > 0$ , that  $\Pi_e^2 \gg \Omega_1 \Omega_e \gg \omega^2$ , and that  $\tan^2 \theta \ll \tan^2 \theta_{\text{Res}}$  ( $\theta_{\text{Res}}$  is the vertex angle of the wave velocity surface).

For the electron densities and magnetic field strengths found in the ionosphere the inequalities  $L > 0$  and  $\Pi_e^2 \gg \Omega_1 \Omega_e \gg \omega^2$  are valid for proton whistlers. The range of wave normal angles for which the inequality  $\tan^2 \theta \ll \tan^2 \theta_{\text{Res}}$  is satisfied can be seen as follows from the equation for the vertex angle of the wave velocity surface: For frequencies near the  $k$ th ion gyrofrequency  $\theta_{\text{Res}}$  is given by

$$\tan^2 \theta_{\text{Res}} = \frac{2}{\alpha_k} (M_k/M_e) \Delta\omega/\Omega_k \quad (5)$$

$$\alpha_k = n(k)/n_e; \Delta\omega = \Omega_k - \omega$$

$$M_k/M_e = \text{Ion Mass/Electron Mass}$$

For proton whistlers  $\Delta\omega/\Omega_1$  is never found to be less than about  $5 \times 10^{-3}$  because of cyclotron damping for frequencies near  $\Omega_1$  [Gurnett and Brice, 1965]. Using this minimum value for  $\Delta\omega/\Omega_1$  it is evident that  $\theta_{\text{Res}}$  is always near  $\pi/2$ . Thus, equation (2) is valid for proton whistler propagation in the ionosphere provided  $\theta$  is not near  $\pi/2$ . Wave normal angles near  $\pi/2$  are not likely because the large refractive index in the ionosphere causes the wave normal to be refracted to near vertical so that  $\theta = \pi/2$  could only occur near the equator. Relatively few proton whistlers are received near the equator [Shawhan, 1965] compared with the occurrence rates at midlatitudes.

In order to further simplify equation (2), we now restrict the discussion to frequencies near an ion gyrofrequency so that  $L \gg R$ . With this assumption the solution for the slow wave from equation (2) is

$$n^2 = L \frac{(1 + \cos^2\theta)}{2 \cos^2\theta} \quad (6)$$

To obtain an expression for  $L(\omega)$  consistent with the assumption that the frequency under discussion is close to an ion gyrofrequency we expand equation (4) about  $\Omega_1$  in terms of  $\Delta\omega = \Omega_1 - \omega$ . Neglecting higher order terms in the parameter  $\Delta\omega/\Omega_1$ , we obtain for the refractive index of ion cyclotron waves near  $\Omega_1$ ,

$$n^2 = \frac{\Pi_1^2}{\Omega_1(\Omega_1 - \omega)} \frac{(1 + \cos^2\theta)}{2 \cos^2\theta} \quad (7)$$

To illustrate the range of frequencies for which this expression is valid the refractive indices for longitudinal propagation,  $n^2 = R$  and  $n^2 = L$ , are plotted in Figure 2. The plasma in this example consists of 80%  $H^+$ , 20%  $He^+$ , and electrons. The approximate expression given by equation (7) is also compared with the exact equation  $n^2 = L$  for  $\theta = 0$ . Equation (7) is seen to be a good approximation for frequencies well above the crossover frequency.

The phase refractive index surface  $n(\theta)$  given by equation (7) is shown in Figure 3. Since the group ray direction is normal to the phase refractive index surface, we see that the ion cyclotron wave energy is constrained to propagate nearly along the static magnetic field. The angle  $\Psi$  between the group ray direction and  $\vec{B}_0$  is given by

$$\tan \Psi = \frac{\sin\theta \cos^3\theta}{1 + \cos^4\theta} \quad (8)$$

From this expression it can be shown that  $\Psi \leq 12.3^\circ$ . Thus, the ray path of proton whistlers near the proton gyrofrequency is very nearly along a geomagnetic field line. For comparison, the ray path of an electron whistler (considering only electron motion) was shown by Storey [1953] to be confined to angles  $\Psi \leq 19.3^\circ$ .

The group velocity is obtained from equation (7) using

$$\left| \frac{\vec{r}}{u} \right| = \frac{c}{n + \frac{\partial n}{\partial \omega}} \left[ 1 + \left( \frac{1}{n} \frac{\partial n}{\partial \theta} \right)^2 \right]^{1/2} \quad (9)$$

We obtain for the component of the group velocity along the static magnetic field ( $u_{||}$ )

$$u_{||} = u_0 \frac{[2(1 + 2 \cos^2 \theta + \cos^6 \theta)]^{1/2}}{(1 + \cos^2 \theta)^{3/2}} \cos \psi \quad (10)$$

$$u_0 = \frac{c \Omega_1^{1/2} (\Omega_1 - \omega)^{3/2}}{\Pi_1 (\Omega_1 - \omega/2)} \quad (11)$$

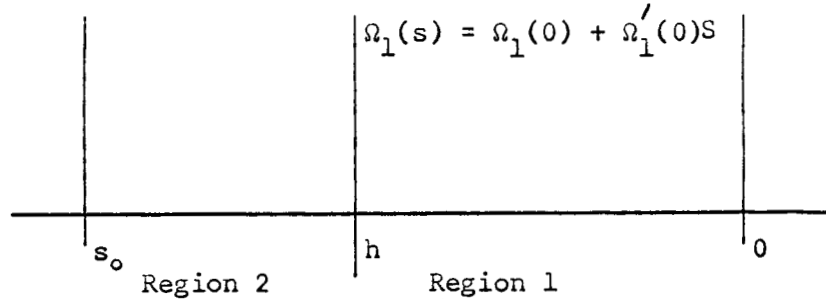
These group velocity expressions are plotted in Figure 3 as a group ray refractive index surface  $\mu(\psi) = c/|\vec{u}|$  and as a plot of  $u_0/u_{||}$  vs  $\theta$ . It is evident from these plots that  $u_{||}$  is nearly independent of  $\theta$  for wave normal angles up to about  $50^\circ$  and is approximated very closely by the group velocity for longitudinal propagation (equation 11).

The group velocity expression for longitudinal propagation  $u_0$  is correct at all frequencies in the limit of a 100%  $H^+$  plasma. For a  $H^+$ ,  $He^+$  plasma the error in using the approximate equation (11) for the group velocity is shown in Figure 4 by comparing equation (11) with the more complicated but exact expression (equation 4). Evidently

this approximate expression is very good for frequencies well above the crossover frequency and rapidly approaches the exact expression as the wave frequency approaches the proton gyrofrequency.

We now consider the travel time integral discussed earlier (equation 1). The ray path is represented schematically as a horizontal axis  $s$  with the wave propagating to the right from an impulse source at  $s_0$ . The static magnetic field increases monotonically to the left with the proton gyrofrequency having a value  $\Omega_1(0)$  at the origin (the observation point). We are only concerned with wave frequencies below and near  $\Omega_1(0)$ .

It is convenient to divide the propagation path into two regions as shown below with the total travel time being written  $t(\omega) = t_1(\omega) + t_2(\omega)$



In Region 1 ( $0 \leq s < h$ ) we assume that the wave is a left hand polarized ion cyclotron wave and that the wave frequency is close to the proton gyrofrequency throughout this region. We also assume that the proton gyrofrequency can be approximated by  $\Omega_1(s) = \Omega_1(0) + \Omega_1'(0)s$ , where

$\Omega_1'(0) = \partial\Omega_1/\partial s$ . In region 2 ( $h < s \leq s_0$ ) we assume that for the frequencies of interest there are no poles or zeros in the refractive index. Thus, the group velocity along this portion of the ray path is always nonzero.\* From this we conclude that  $t_2(\omega)$  will be a finite continuous function of  $\omega$  for the frequencies of interest (since  $u \neq 0$  anywhere along the path).

In the application of this propagation model to proton whistler propagation in the ionosphere,  $s_0$  is at the base of the ionosphere ( $\approx 100$  km),  $s = 0$  is the position of the satellite, and the polarization reversal which gives rise to proton whistlers occurs in Region 2.

Since the wave in Region 1 is an ion cyclotron wave with  $\omega$  near  $\Omega_1$  we know that the ray direction is nearly along the static magnetic field. This allows us to use equation 11 for  $u$  in the travel time integral  $t_1(\omega)$ :

$$t_1(\omega) = \frac{1}{c} \int_h^0 \frac{\Pi_1(s)[\Omega_1(s) - \omega/2]ds}{\Omega_1(s)^{1/2}[\Omega_1(s) - \omega]^{3/2}} \quad (12)$$

For small  $\Delta\omega = \Omega_1(0) - \omega$  the kernel multiplying  $\Pi_1(s)$  in the integrand is large only in the neighborhood of  $s = 0$ . We therefore assume that  $\Pi_1(s)$  varies slowly in Region 1 so that we may write: (using  $\Omega(s) = \Omega_1(0) + \Omega_1'(0)s$ )

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\*This is equivalent to saying that  $h$  is accessible from  $s_0$ .

$$t_1(\omega) = \frac{\Pi_1(0)}{c} \int_h^0 \frac{(\Omega_1(0) + \Omega_1'(0)s - \frac{\omega}{2})ds}{[\Omega_1(0) + \Omega_1'(0)s]^{1/2} [\Omega_1(0) - \omega + \Omega_1'(0)s]^{3/2}} \quad (13)$$

we obtain for the integral

$$t_1(\omega) = \frac{\Pi_1(0)}{c \Omega_1'(0)} \left[ \frac{1}{[1 - \frac{\omega}{\Omega_1(0)}]^{1/2}} - \frac{1}{[1 - \frac{\omega}{\Omega_1(h)}]^{1/2}} + 2 \sinh^{-1} \frac{[\frac{\Omega_1(0)}{\omega} - 1]}{\omega} - 2 \sinh^{-1} \frac{[\frac{\Omega_1(h)}{\omega} - 1]}{\omega} \right] \quad (14)$$

Expanding to the first order in terms of the small quantity

$\Delta\omega/\Omega_1$  we obtain:

$$t_1(\omega) = \frac{\Pi_1(0)\Omega_1(0)^{1/2}}{c \Omega_1'(0)} \left[ \frac{1}{\Delta\omega} \right]^{1/2} + [\text{Terms} \propto (\Delta\omega)^n] \quad (15)$$

We note that  $h$  occurs only in the higher order terms in  $\Delta\omega$ ; hence, in the limit  $\Delta\omega \rightarrow 0$  the dependence of  $t_1(\omega)$  on  $h$  appears only as a constant.

Since  $t_2(\omega)$  is finite and continuous about  $\Omega_1$  we have for small  $\Delta\omega$ :  $t_2(\omega) = t_2(\Omega_1) + t_2'(\Omega_1) \Delta\omega$ . The asymptotic expression for  $t(\omega)$  in the limit  $\Delta\omega \rightarrow 0$  is dominated by  $t_1(\omega)$

and therefore is

$$t(\omega) = \frac{\Pi_1(0)\Omega_1(0)^{1/2}}{c \Omega_1'(0)} \left[ \frac{1}{\Delta\omega} \right]^{1/2} + \text{constant} \quad (16)$$

This equation gives the asymptotic form for the long drawn out "high frequency tail" of the proton whistler for  $\omega$  near the proton gyrofrequency at the satellite (see Figure 1). This relatively simple relation, independent of the concentrations along all but the latter portion of the ray path, occurs because of the singular nature of the travel time integral near the proton gyrofrequency.

Equation (16) is the working equation for the determination of  $n(H^+)$  and  $\Omega_1$  from proton whistlers. The interpretation to be given to equation (16) is as follows:

Assume that the local gyrofrequency at the satellite is  $\Omega_1$ , then the travel time  $t(\omega)$  for the portion of proton whistler signal near the proton gyrofrequency is proportional to the square root of the period of the difference frequency  $\Delta\omega = \Omega_1 - \omega$ . If  $t(\omega)$  vs  $(1/\Delta\omega)^{1/2}$  is not found to be a straight line then the assumed gyrofrequency is incorrect and must be adjusted until the best straight line fit is obtained.



This gyrofrequency is then a determination of the true proton gyrofrequency at the satellite.

Since  $\partial\Omega_1/\partial s$  can be accurately calculated at the position of the satellite from the geometry of the geomagnetic field, it is evident from equation (16) that the slope of  $t(\omega)$  vs  $(1/\Delta\omega)^{1/2}$  can be used to obtain  $\Pi_1(0)$  and hence  $n(H^+)$  from

$$\Pi_1^2(0) = \frac{4\pi e^2 n(H^+)}{m_1}$$

The experimental method used to calculate  $\Pi_1(0)$  and  $\Omega_1(0)$  from proton whistlers and the accuracy of the method are discussed in the next section.

### III. THE METHOD USED TO DETERMINE $n(H^+)$ AND $\Omega_1(0)$ FROM THE DISPERSION OF PROTON WHISTLERS

We have shown that equation (16) gives the travel time of a proton whistler in the limit of small  $\Delta\omega$ . We have also suggested that equation (16) be used to determine  $n(H^+)$  and  $\Omega_1(0)$  from the dispersion of proton whistlers received by a satellite. In the following we discuss a method of calculating  $n(H^+)$  and  $\Omega_1(0)$  from a set of measured proton whistler travel times at different frequencies ( $t_j, \omega_j$ ). Since actual measurements must be made for a finite range of  $\Delta\omega$  values and since equation (16) is strictly valid only in the limit  $\Delta\omega \rightarrow 0$  we must estimate the error in calculating  $n(H^+)$  and  $\Omega_1(0)$  from equation (16) using a realistic range of  $\Delta\omega$  values. This we do by testing the accuracy of equation (16) using as data calculated proton whistler travel times for an assumed model ionosphere. By comparing the values of  $n(H^+)$  and  $\Omega_1(0)$  for the assumed ionospheric model with the values obtained from fitting equation (16) to the calculated proton whistler travel times we can estimate the errors which arise from using equation (16) for finite values of  $\Delta\omega$  (independent of measurement errors). The following is an illustration of one such test case.

In Figure 5 is shown one of several model ionospheres used to test equation (16). The ion density profiles for this

model are calculated using diffusive equilibrium relations with an ion temperature of 800°K. The detailed choice of the model ionosphere is not crucial since we only wish to illustrate the method of determining  $n(H^+)$  and  $\Omega_1(0)$  and to estimate the errors in the method.

Using the ionospheric model of Figure 5 proton whistler travel times  $t(\omega)$  to an altitude of 2000 km are calculated and shown in Figure 1. These travel times are calculated using the exact equations for longitudinal propagation as discussed by Gurnett et al. [1965]. In Figure 6 we plot  $t(\omega)$  vs  $[1/\Delta\omega^*]^{1/2}$ ,  $\Delta\omega^* = \Omega_1^* - \omega$ , for this test case, where  $\Omega_1^*$  is a trial value for the proton gyrofrequency. Five curves are shown for  $\Omega_1^*$  values differing by 1 cps. The actual proton gyrofrequency at 2000 km for this test case is  $\Omega_1(0) = 307.3$  cps. The relation between  $t(\omega)$  and  $[1/\Delta\omega^*]^{1/2}$  is seen to be very nearly linear when  $\Omega_1^* = 307$  cps  $\approx \Omega_1(0)$ . However, when  $\Omega_1^*$  is as little as 1 cps different from  $\Omega_1(0)$  the curves deviate markedly from a straight line.

The plots in Figure 6 show that  $\Omega_1(0)$  can be determined from proton whistler measurements by finding the value of  $\Omega_1^*$  for which the measured pairs of points  $(t_j, p_j = [1/(\Omega_1^* - \omega_j)]^{1/2})$  lie on a straight line. This value, call it  $\hat{\Omega}_1$ , is the best estimate of the proton gyrofrequency at the satellite. The  $H^+$

concentration is determined from the slope  $S = \Delta t / \Delta p$  (evaluated at  $\Omega_1^* = \hat{\Omega}_1$ ) using the following relation derived from equation (16):

$$n(H^+) = \frac{m_1 c^2}{4\pi e^2} \frac{(\partial \Omega_1 / \partial S)^2}{\hat{\Omega}_1(0)} S^2 \quad (18)$$

To implement this scheme for calculating  $\Omega_1(0)$  and  $n(H^+)$  we use the statistic  $T(\Omega_1^*)$  to find the value of  $\Omega_1^*$  which gives the best fitting linear relation between the points  $(t_j, p_j)$ .

$$T(\Omega_1^*) = \frac{\sum_j^n (t_j - \bar{t})(p_j - \bar{p})}{\left[ \frac{\sum_j (p_j - \bar{p})^2 \sum_j (t_j - \bar{t})^2 - [\sum_j (t_j - \bar{t})(p_j - \bar{p})]^2}{n-2} \right]^{1/2}} \quad (19)$$

$$\bar{t} = \frac{1}{n} \sum t_j, \quad \bar{p} = \frac{1}{n} \sum p_j$$

The value of  $T$  is a maximum when the mean square error between the experimental points and the best fit straight line is a minimum [Lacey, 1959]. When the experimental points lie exactly in a straight line it is easy to show that the statistical parameter  $T$  is infinite.

In Figure 7 is shown a plot of  $T(\Omega_1^*)$  vs  $\Omega_1^*$  for the calculated proton whistler shown in Figure 1. The points

$(t_j, \omega_j)$  used to calculate  $T(\Omega_1^*)$  are 1 cps apart in frequency for  $0 < \Delta\omega < 40$  cps. To prevent the pole in  $p_j$  at  $\Delta\omega^* = 0$  from dominating the sums in equation (19) the points  $(t_j, p_j)$  are included in the sums only if  $\Delta\omega^* > 1.0$  cps.<sup>†</sup> This condition gives rise to small discontinuities in  $T(\Omega_1^*)$  as new points are added to the sums in equation (19) (see Figure 7). The sharp peak in  $T(\Omega_1^*)$  at  $\Omega_1^* = 370.06$  cps indicates that the straight line fit through the points  $(t_j, p_j)$  is the best for this value of  $\Omega_1^*$ . This conclusion is consistent with the curves shown in Figure 6 which show that the best straight line fit occurs when  $\Omega_1^*$  is approximately 370 cps. From the position of the peak in  $T$  we estimate the proton gyrofrequency at 2000 km to be 370.06 cps. This gyrofrequency is in error by 0.24 cps (0.05%) from the gyrofrequency used in the model ionosphere. This small error arises because we have used a finite range of  $\Delta\omega$  values (40 cps) in equation (16) whereas this equation is strictly valid only in the limit  $\Delta\omega \rightarrow 0$ . This error is unavoidable in the sense that actual measurements must be made over a finite range of  $\Delta\omega$  values.

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<sup>†</sup> This requirement is justifiable because cyclotron damping strongly attenuates frequencies for which  $\Delta\omega$  is less than about 2 cps.

To calculate  $n(H^+)$  from equation (18) the slope of the best fit straight line through the points  $(t_j, p_j)$  is calculated using

$$S(\Omega_1^*) = \frac{\Sigma(t_j - \bar{t})(p_j - \bar{p})}{\Sigma(p_j - \bar{p})^2} \quad (20)$$

In Figure 7 is shown a plot of  $n(H^+)$  vs  $\Omega_1^*$  for the calculated proton whistler shown in Figure 1. From the peak in  $T(\Omega_1^*)$  the  $H^+$  concentration is estimated to be  $3.62 \times 10^3 \text{ (cm}^{-3}\text{)}$ . The actual  $H^+$  concentration used in the model ionosphere at 2000 km altitude is  $3.76 \times 10^3 \text{ (cm}^{-3}\text{)}$ , an error of 3.8%.

Using the ionospheric model shown in Figure 5 test cases similar to the one just described have been analyzed for different altitudes. The error in calculating  $n(H^+)$  and  $\Omega_1(0)$  for these cases is summarized in Table 1. For all cases the measured frequencies are in the range  $0 < \Delta\omega < 40$  cps. Table 1 thus gives an estimate of the errors in determining  $\Omega_1(0)$  and  $n(H^+)$  from equation (16) when there are no measurement errors. For actual proton whistler data the peak in  $T(\Omega_1^*)$  is not nearly as sharp as for these test cases because of experimental errors in measuring the points  $(t_j, \omega_j)$ . To determine error bars we calculate  $n(H^+)$  and  $\Omega_1(0)$  at the peak

of  $T(\Omega_1^*)$  for several independently measured data sets  $(t_j, \omega_j)$  from the same proton whistler. Error bars are given by the standard deviation of these repeated measurements.

Another possible method of fitting equation (16) to observed proton whistler signals is based on maximizing the correlation coefficient between an ideal signal as predicted by equation (16) and the actual proton whistler signal. This method is an extension of a method used by Beghin and Siredy [1964] to analyze electron whistler travel times. The correlation coefficient is of the following form (suitably normalized)

$$\rho[t_0, n(H^+), \Omega_1^*] = A \int_{-\infty}^{\infty} S(t) \cos \Phi(t) dt$$

$$\Phi(t) = \int \omega(t + t_0) dt$$

The function  $S(t)$  is the received proton whistler signal and  $\omega(t)$  is obtained by solving equation (16) for  $\omega$  as a function of  $t$ . The correlation coefficient is a function of three unknown quantities: the time origin  $t_0$ , the  $H^+$  concentration, and the proton gyrofrequency. Values for these unknown quantities can be determined by finding the value for these three quantities which maximizes the correlation coefficient. The digitizing and computer programming involved in this method requires a considerable investment in equipment and time so we have not pursued this method.

#### IV. MEASUREMENTS OF $\Omega_1(0)$ AND $n(H^+)$ FROM PROTON WHISTLERS RECEIVED BY INJUN 3

Up to this time we have assumed that some suitably accurate method exists to measure the travel time of a proton whistler as a function of frequency. Two methods are used. First, points  $(t_j, \omega_j)$  can be measured directly from spectrograms of proton whistlers and second, the points  $(t_j, \omega_j)$  can be measured from oscillograph records of the difference frequency signal derived by beating the proton whistler signal against a fixed frequency oscillator. The records obtained using these two methods are contrasted in Figure 8 for a proton whistler received with Injun 3. The spectrogram in this illustration was made on a Sonograph frequency-time spectrum analyzer (Kay Electric, Pine Brook, New Jersey). The oscillograph record shown gives the difference frequency between the proton whistler and a fixed frequency oscillator adjusted to a frequency  $\omega_{osc}$  near the proton gyro-frequency (527.5 cps for the case illustrated). A block diagram of the equipment used to obtain this oscillograph record is shown in Figure 9. A stable 10 kc/sec reference tone recorded on the data tape provides the time base for accurately determining the oscillator frequency and providing time markers on the oscillograph record. The band width of the low pass filter is usually set to about 15 cps. The proton whistler frequency at time  $t_j$  is



calculated using  $\omega_j = \omega_{osc} - 1/T_j$  (cps), where  $T_j$  is the period of the beat frequency signal at time  $t_j$ . Usually measurements from both the sonograph record and the beat frequency record are used. These measurements are complementary in that the beat frequency records give excellent accuracy for wave frequencies near the proton gyrofrequency ( $\Delta\omega$  less than about 15 cps); whereas, the sonograph records give good accuracy for larger values of  $\Delta\omega$  (greater than about 15 cps).

Nineteen points  $(\omega_j, t_j)$  have been measured for the proton whistler shown in Figure 8. In Figure 10 is shown a plot of  $[1/(\Omega_1^* - \omega)]^{1/2}$  vs  $t$  for five values of  $\Omega_1^*$  as calculated from the  $(\omega_j, t_j)$  measurements. In Figure 11 is shown a plot of the statistical parameter  $T(\Omega_1^*)$  for these measurements. The peak in the parameter  $T(\Omega_1^*)$  occurs when  $\Omega_1^* = 528.3$  cps and  $n(H^+) = 1.95 \times 10^3$  ( $\text{cm}^{-3}$ ). These values are the best estimate of  $\Omega_1(0)$  and  $n(H^+)$  at the satellite from measurements of the oscillograph record in Figure 8. By independently analyzing five independent records for this same proton whistler signal (not all having the same oscillator frequency) the following error bars (one standard deviation) due to random measurement errors have been determined for this proton whistler:

$$\begin{aligned}\Omega_1(0) &= 528.0 \pm 0.9 \text{ (cps)} \\ n(H^+) &= (1.90 \pm 0.28) \times 10^3 \text{ (cm}^{-3}\text{)}\end{aligned}$$

From the crossover frequency, estimated to be within  $278 < \omega_{12}$

$< 330$  (cps), the fractional concentration  $\mathcal{V}_1 = n(H^+)/n_e$  has been calculated to be [see Shawhan and Gurnett, 1965]  $.62 < \mathcal{V}_1 < .74$  for this proton whistler. From  $\mathcal{V}_1$  we obtain the following estimate of the electron number density at the satellite,

$$n_e = (2.97 \pm 0.50) \times 10^3 \text{ (cm}^{-3}\text{)}.$$

In Table 2 we present a number of values of  $\Omega_1(0)$ ,  $n(H^+)$ , and  $n_e$  calculated from proton whistlers received by Injun 3. Also shown for comparison are  $\Omega_1(0)$  values calculated from the Jensen and Cain expansion for the geomagnetic field at the satellite [Jensen and Cain, 1962] and  $n_e$  values obtained from the Alouette 1 satellite for the same latitude, season, and local time as the Injun 3 measurements. The Alouette electron density measurements were kindly provided by M. J. Rycroft of Ames Research Center. The largest disagreement between the proton gyrofrequency determined from proton whistlers and from the Jensen and Cain expansion is 1%. This disagreement we attribute to the accuracy limitations of the Jensen and Cain expansion for the geomagnetic field. The electron densities calculated from proton whistlers are seen to be in reasonably good agreement with electron densities obtained from the Alouette 1 satellite at 1000 km. This general agreement with other data is interpreted as establishing the validity of the proton whistler method as another independent method for the determination of magnetic fields ( $\Omega_1(0)$ ),  $n(H^+)$ , and  $n_e$  in the ionosphere.

## V. CONCLUSION

In this paper the propagation of ion cyclotron waves near an ion gyrofrequency was discussed. It was found that near an ion gyrofrequency the ray path of the left-hand polarized wave is very nearly along the static magnetic field line and that the parallel component of the group velocity is nearly independent of wave normal angle up to about  $50^\circ$ . A simple equation for the parallel component of the group velocity was given.

These equations were applied to the travel time integral of proton whistlers from the source of the lightning impulse to the satellite. It was found that for frequencies near the proton gyrofrequency at the satellite the kernel multiplying  $n(H^+)^{1/2}$  in the integrand is large only in the region near the satellite. Assuming  $n(H^+)$  is approximately uniform within this region an expression was derived for the frequency-time trace of a proton whistler near the proton gyrofrequency. It was shown that the proton gyrofrequency and the  $H^+$  concentration in the vicinity of the satellite could be obtained by fitting this theoretical frequency-time expression to observed proton whistler signals. By also measuring the crossover frequency of the proton whistler the electron density could be calculated.

This method for determining proton gyrofrequency,  $H^+$  concentration, and electron density was applied to proton whistlers received by the Injun 3 satellite. The calculated values were shown to be in good agreement with measurements by other experimenters at similar altitudes, latitudes, and local times.

This new technique for measuring electron densities is expected to be particularly useful in the VLF studies with the Injun 3 satellite because no other detector was included to determine  $n_e$ . This technique illustrates that a considerable amount of information concerning the composition of the ionosphere can be obtained from a VLF receiver of the type flown on Injun 3. Should  $He^+$  and  $O^+$  whistlers be detected by future VLF experiments we expect that a technique similar to the one discussed in this paper could be applied to these whistlers to obtain the  $He^+$  and  $O^+$  number densities.

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## FIGURE CAPTIONS

- Figure 1. Calculated proton whistler travel time for the model ionosphere shown in Figure 5.
- Figure 2. Refractive index for longitudinal propagation compared with the approximate expression given by equation (7).
- Figure 3. Phase and group refractive index surfaces for ion cyclotron waves and  $u_o/u_{||}$  as a function of wave normal angle.
- Figure 4. Group velocity for longitudinal propagation for a  $H^+$ ,  $He^+$  plasma with  $\alpha_1 = 0.8, 0.6$ , and  $0.4$  as compared with the approximate expression given by equation (11).
- Figure 5. An ionospheric model used to calculate proton whistler travel times for testing the accuracy of equation (16) when  $\Delta\omega$  is finite.
- Figure 6. A plot of  $t(\omega)$  vs  $[1/\Delta\omega]^{1/2}$  for the calculated proton whistler travel time shown in Figure 1.
- Figure 7. A plot of  $T(\Omega_1^*)$  and  $n(H^+)$  as a function of the assumed proton gyrofrequency  $\Omega_1^*$  for the calculated proton whistler travel time shown in Figure 1.
- Figure 8. A sonograph record of a proton whistler signal received by Injun 3 and a beat frequency oscillograph record for the same whistler.

Figure 9. A block diagram of the equipment used to derive the beat frequency record shown in Figure 8.

Figure 10. A plot of  $1/(\Omega_1^* - \omega)^{1/2}$  vs  $t$  for the proton whistler shown in Figure 8.

Figure 11. A plot of  $T(\Omega_1^*)$  and  $n(H^+)$  as a function of the assumed proton gyrofrequency  $\Omega_1^*$  for the proton whistler shown in Figure 8.



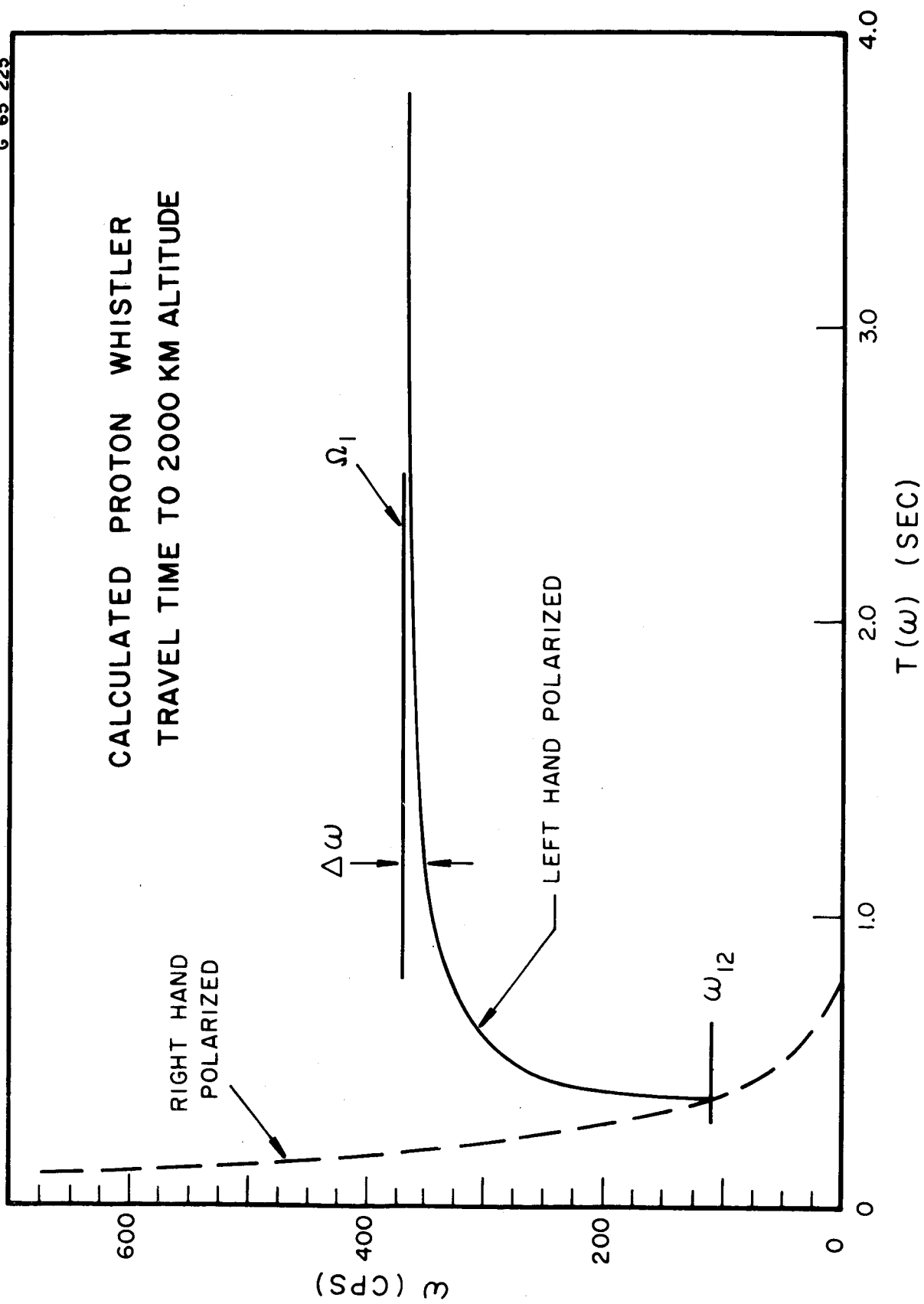


Figure 1

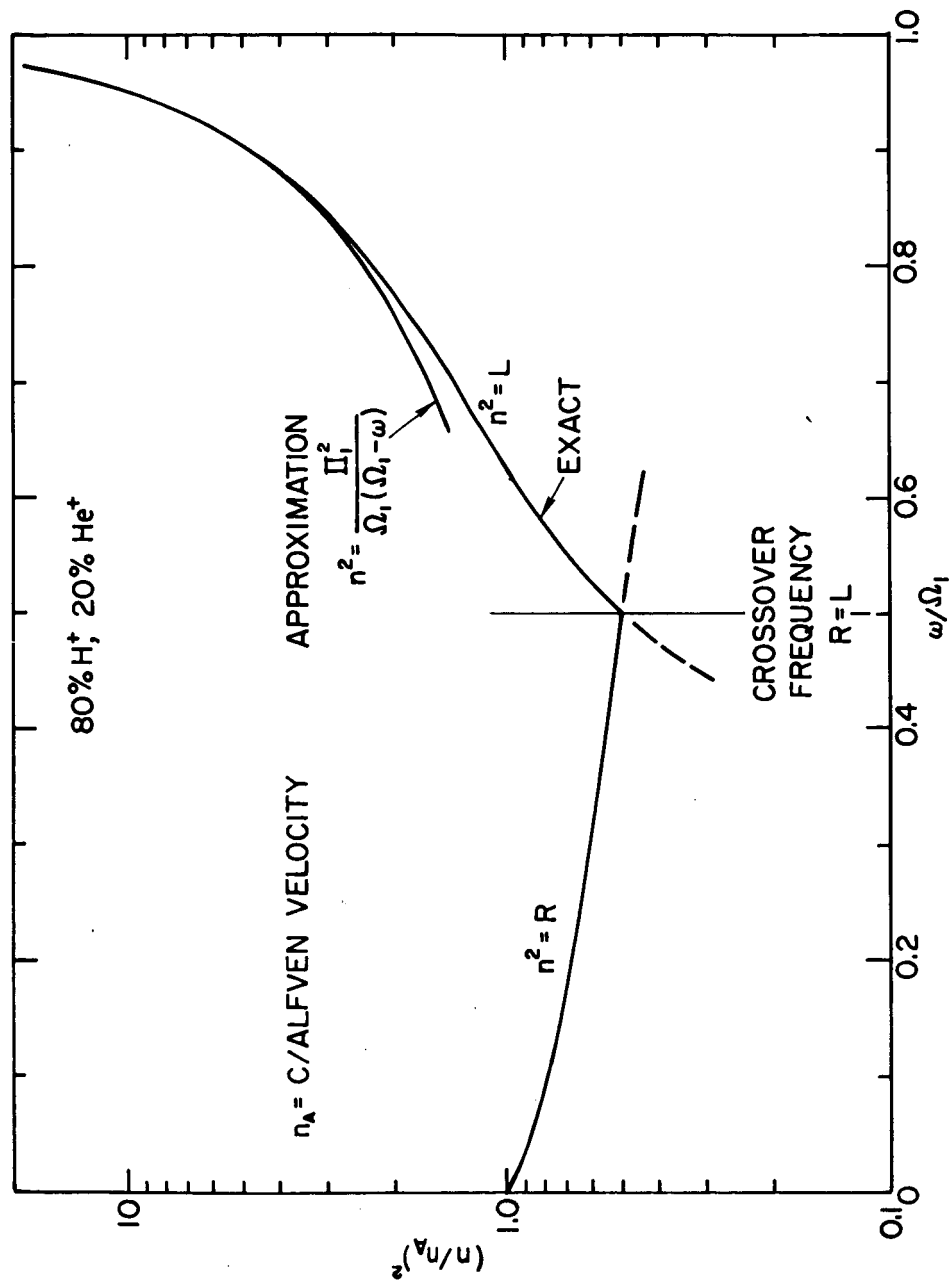


Figure 2

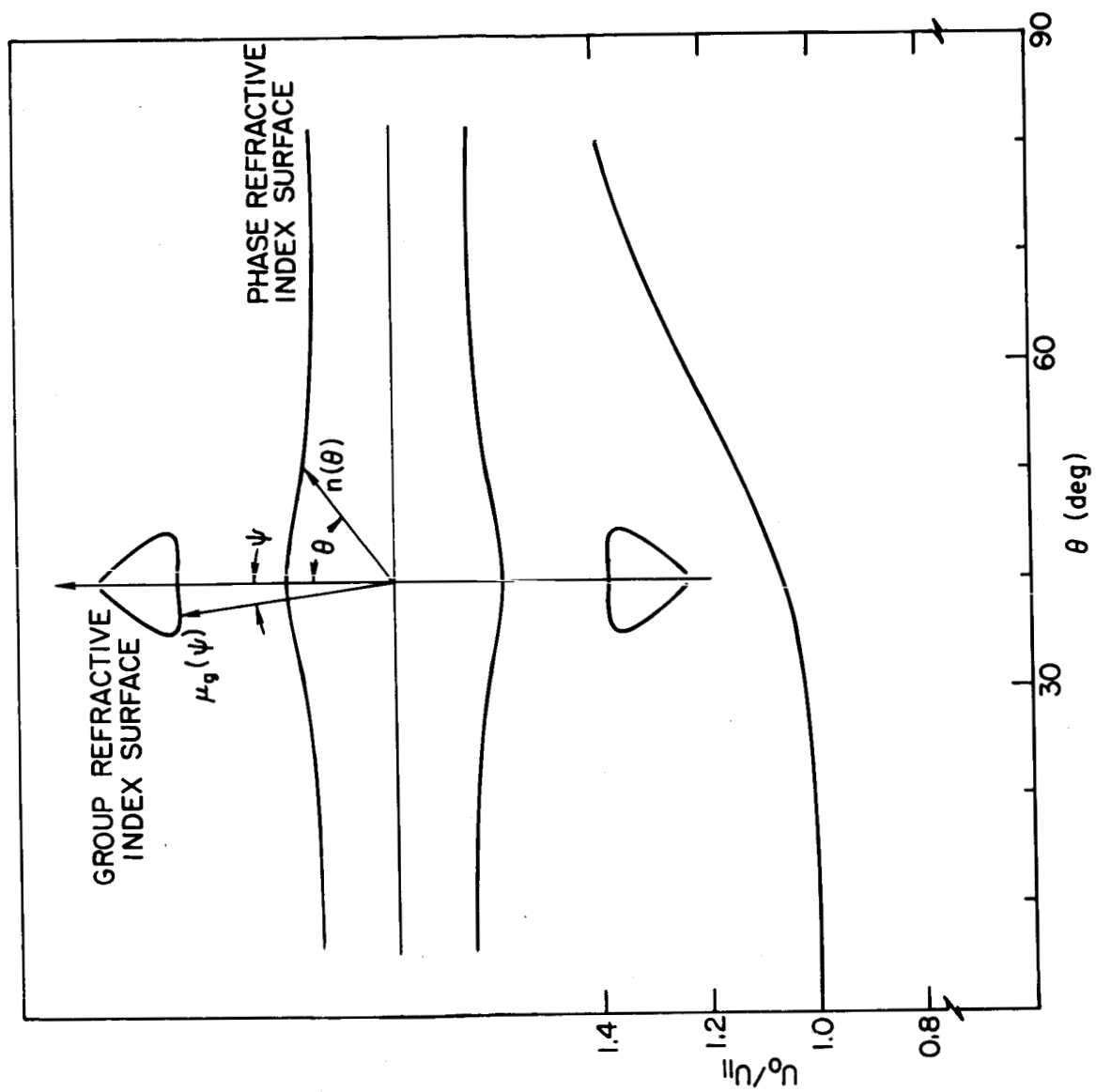


Figure 3

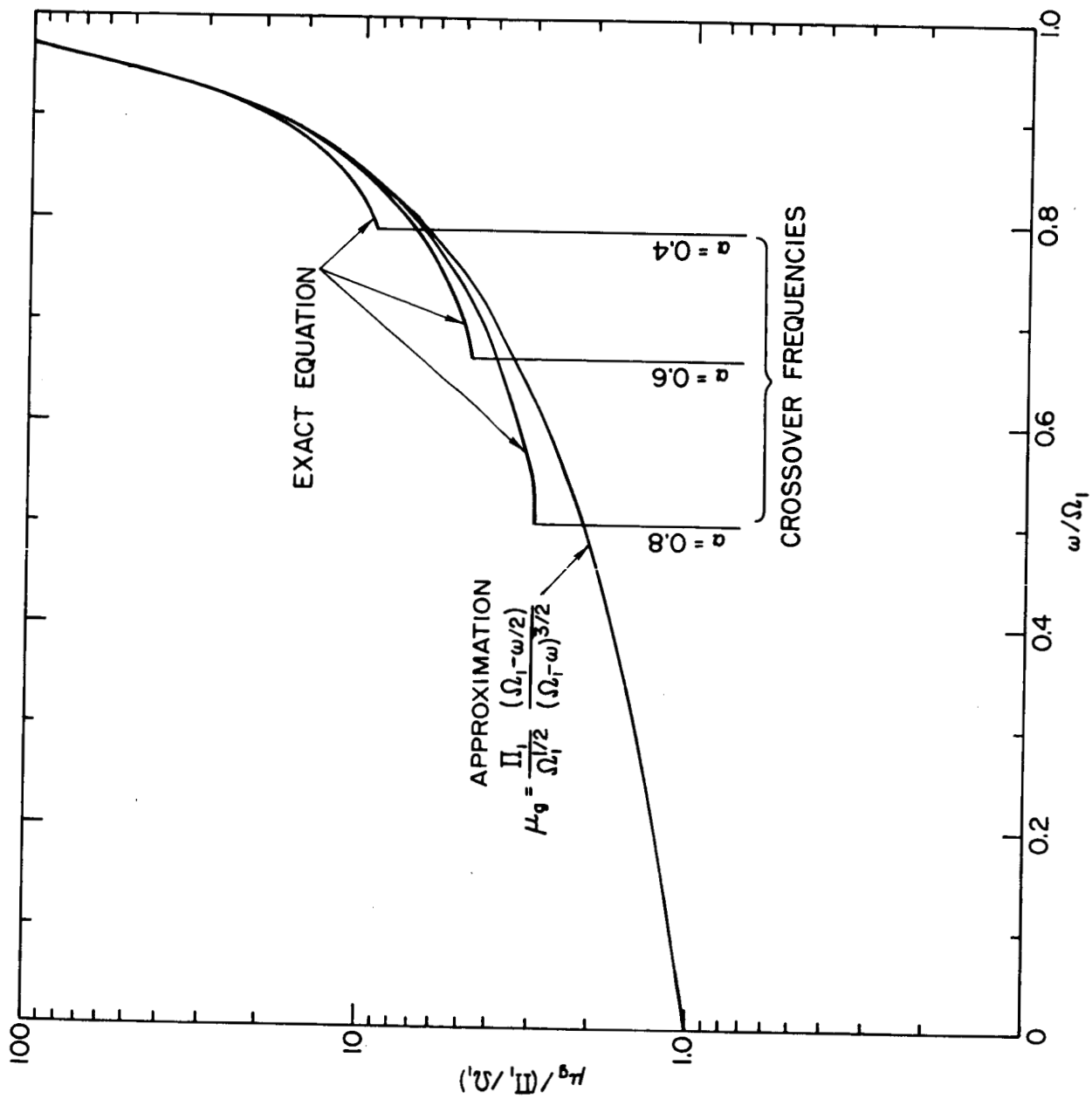
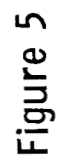


Figure 4



## Figure 5

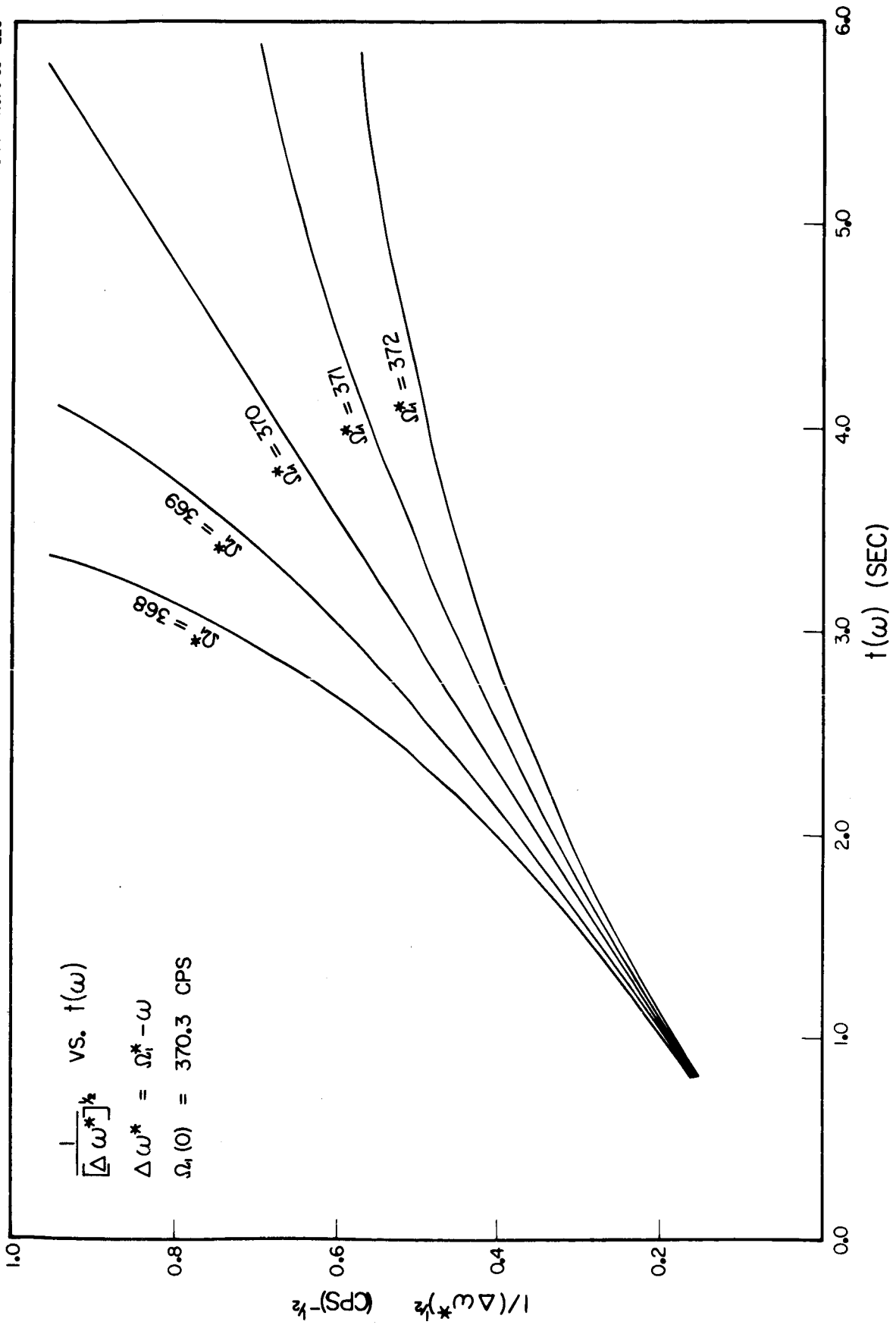


Figure 6

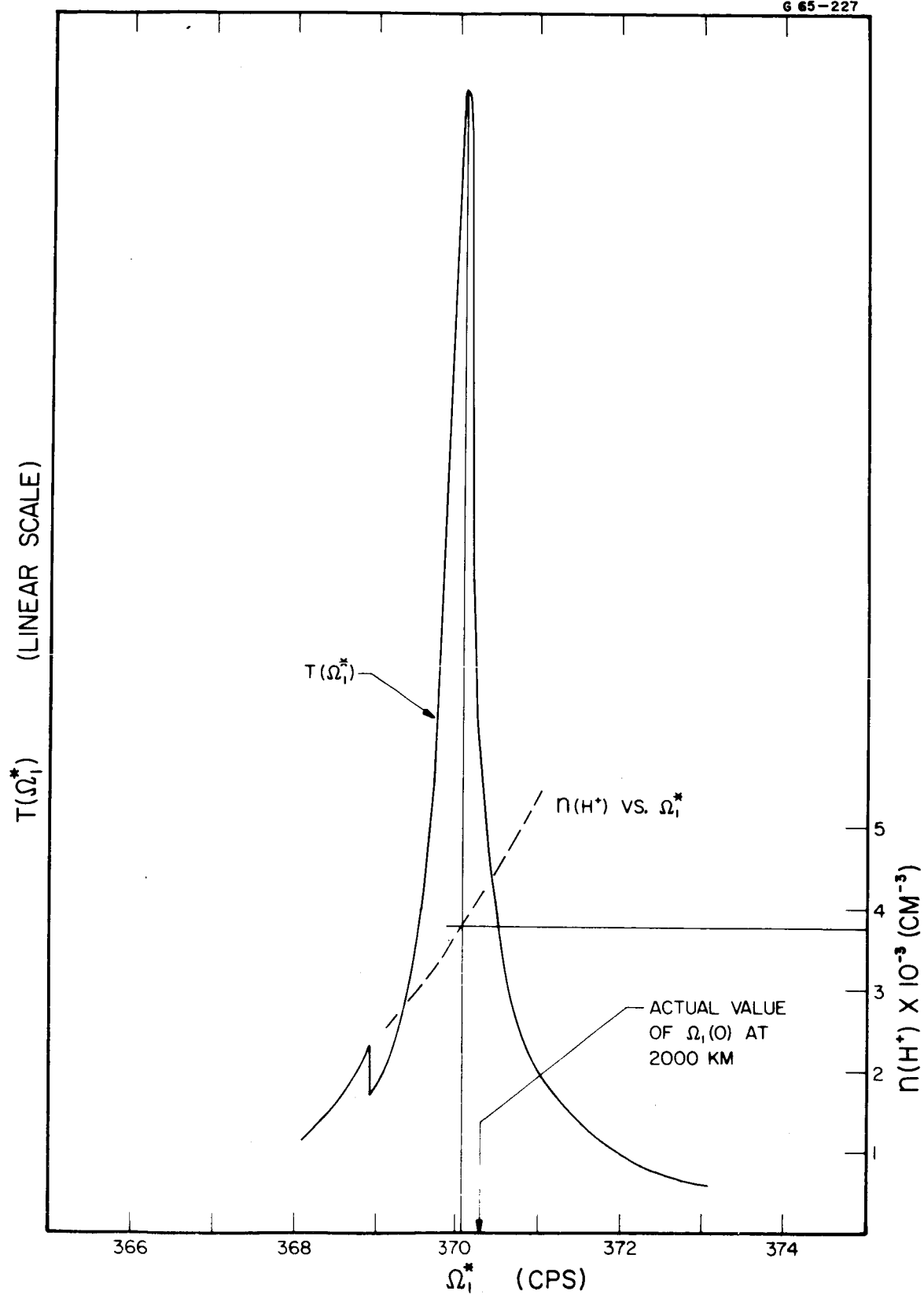


Figure 7

INJUN III PROTON WHISTLER JAN 11, 1963 09:08:26 UT

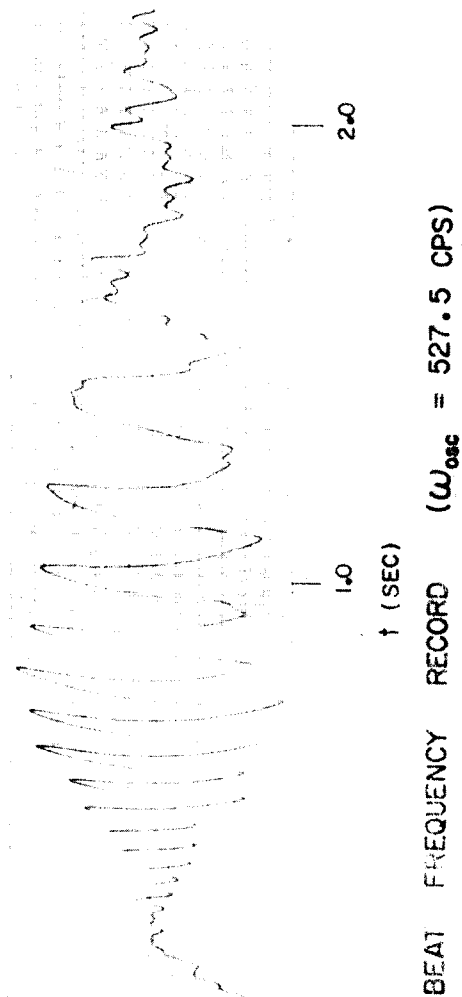
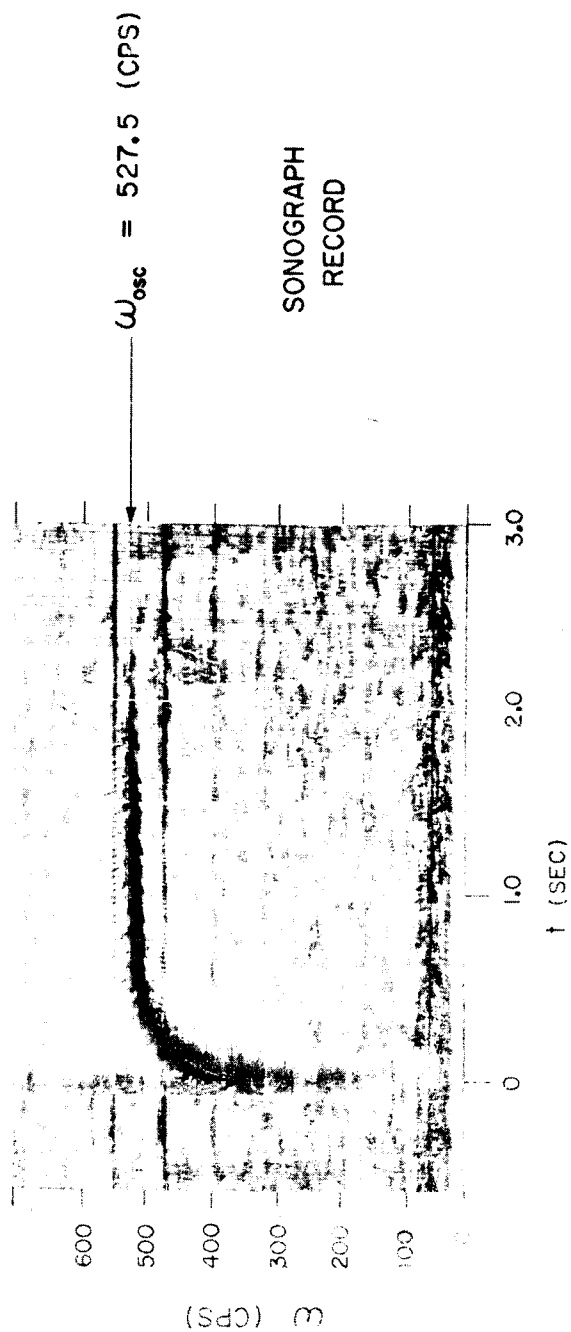


Figure 8



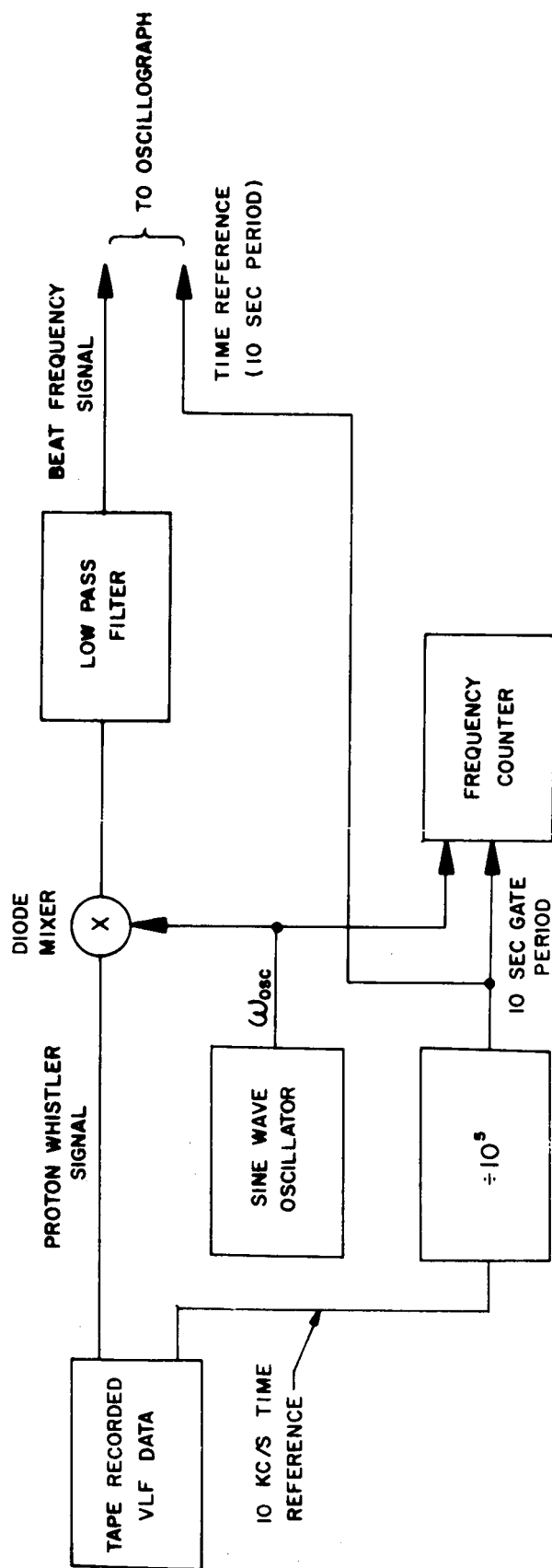


Figure 9

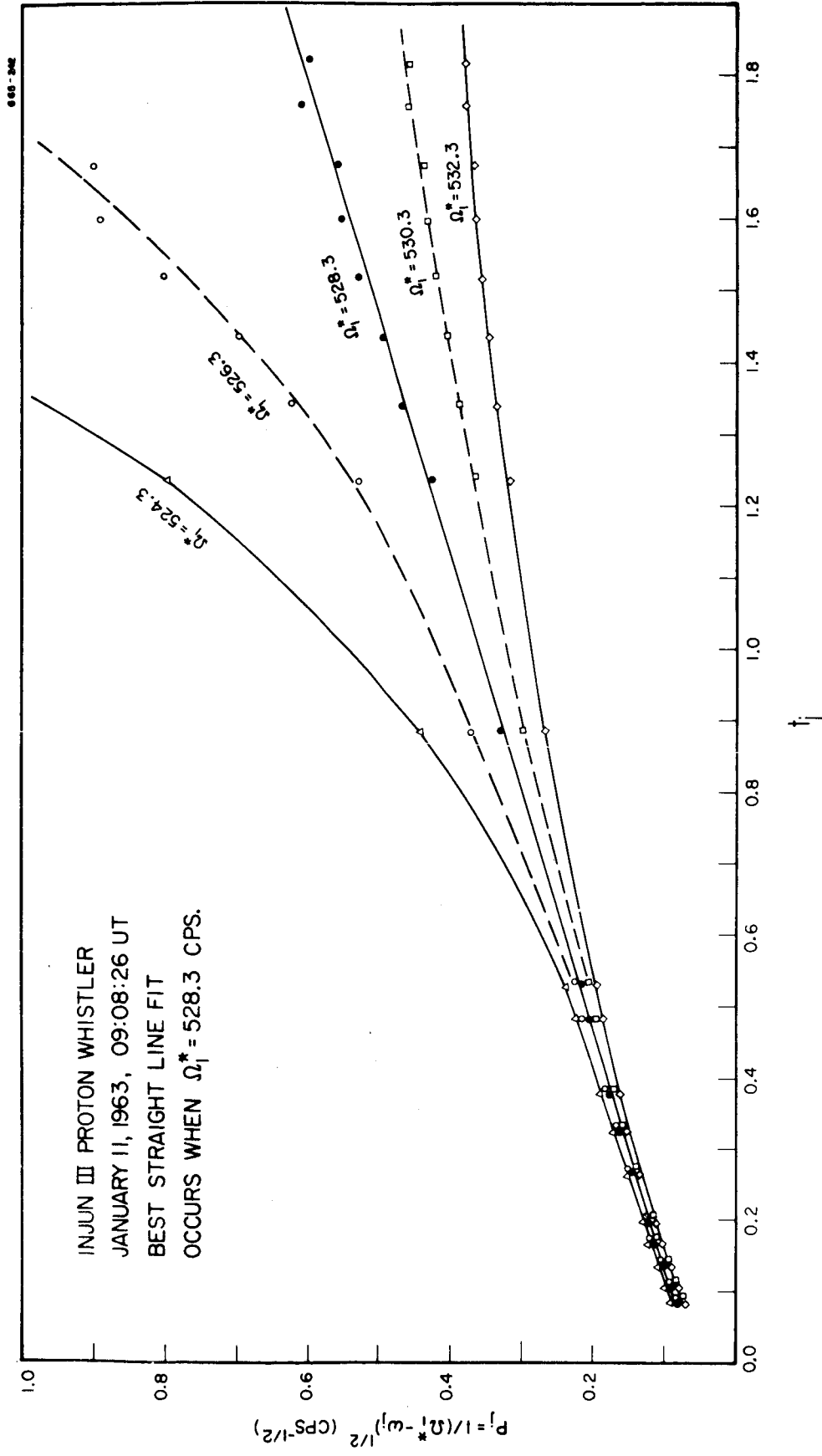


Figure 10

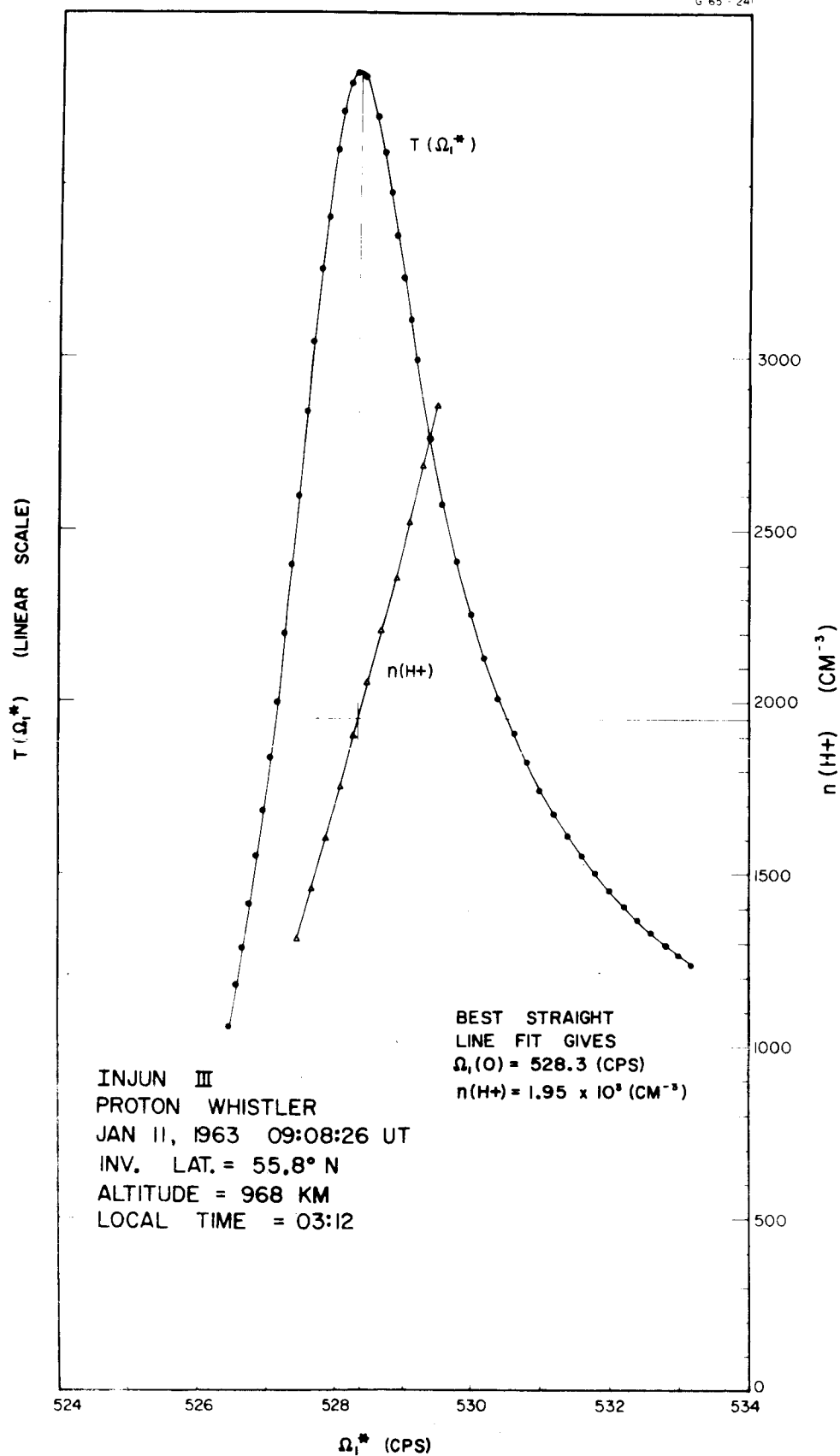


Figure 11

TABLE 1

Altitude (Km)	$\alpha_1$	Model $\Omega_1(0)$ cps	Calculated $\hat{\Omega}_1(0)$ cps	Error %	Model $n(H^+)$ $cm^{-3}$	Calculated $n(H^+)$ $cm^{-3}$	Error %
800	0.34	589.5	588.3	0.17	$4.25 \times 10^3$	$3.42 \times 10^3$	24
1000	0.54	542.6	541.8	0.15	$4.77 \times 10^3$	$4.24 \times 10^3$	11
1200	0.70	500.7	500.2	0.10	$4.82 \times 10^3$	$4.57 \times 10^3$	5
1600	0.88	429.0	428.6	0.09	$4.33 \times 10^3$	$4.30 \times 10^3$	1
200	0.94	370.3	370.06	0.05	$3.76 \times 10^3$	$3.62 \times 10^3$	4

TABLE 2

SATELLITE COORDINATES					PROTON WHISTLER DETERMINATIONS			Nearest Alouette Electron Density Data at 1000 Km
WN = Winter Night, SN = Summer Night					$\Omega_1(0)$	$n(H^+) \times 10^{-3}$	$n_e \times 10^{-3}$	
Universal Time		Altitude	Invariant Latitude	Jensen & Cain Value for $\Omega_1(0)$	cps	$cm^{-3}$	$cm^{-3}$	$n_e \times 10^{-3}$
1963		Km	deg.					
Jan 7 10:08:19	WN	1595	40.2	329.2	$329.9 \pm 0.5$	$7.93 \pm 0.85$	$9.51 \pm 1.00$	
Jan 7 10:13:22	WN	1249	50.2	450.0	$451.5 \pm 0.8$	$7.72 \pm 1.50$	$13.5 \pm 2.5$	7.5 - 12.5
Jan 11 09:08:26	WN	968	55.8	528.2	$528.0 \pm 0.9$	$1.90 \pm 0.28$	$2.97 \pm 0.50$	3.0 - 7.5
Jan 11 10:59:38	WN	1308	42.4	408.5	$412.3 \pm 0.7$	$14.0 \pm 2.6$	$20.4 \pm 3.8$	10.0 - 20.0
Jan 11 11:03:12	WN	1058	51.7	516.4	$517.9 \pm 1.1$	$5.94 \pm 2.58$	$7.46 \pm 3.21$	7.5 - 12.5
Mar 3 07:46:06	WN	1708	48.9	369.0	$373.0 \pm 0.6$	$4.80 \pm 0.96$	$5.93 \pm 1.05$	
June 11 01:42:54	SN	1241	41.1	382.0	$382.4 \pm 0.3$	$5.46 \pm 0.77$	$9.80 \pm 1.50$	7.5 - 12.5
June 11 01:47:06	SN	1535	52.1	387.3	$387.5 \pm 0.6$	$6.09 \pm 1.13$	$8.01 \pm 1.48$	

Table 2

UNCLASSIFIED

Security Classification

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		2b. GROUP	
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4. DESCRIPTIVE NOTES (Type of report and inclusive dates) Progress			
5. AUTHOR(S) (Last name, first name, initial) Gurnett, Donald A., and Shawhan, Stanley D.			
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10. AVAILABILITY/LIMITATION NOTICES Qualified requesters may obtain copies of this report from DDC.			
11. SUPPLEMENTARY NOTES		12. SPONSORING MILITARY ACTIVITY Office of Naval Research	
13. ABSTRACT <p>In this paper we discuss a method for determining <math>H^+</math> concentration, electron number density, and proton gyrofrequency in the vicinity of the satellite by measurements of the asymptotic frequency-time profile of a proton whistler near the proton gyrofrequency. This new technique is applied to proton whistlers received by the Injun 3 VLF receiver. The calculated values of <math>H^+</math> concentration and electron density are shown to be in good agreement with measurements by other experimenters at similar altitudes, latitudes, and local times. B values calculated from the proton gyrofrequency are compared with values calculated from the Jensen and Cain expansion for the geomagnetic field.</p> <p>It is shown that the wave energy of a proton whistler is guided very nearly along the geomagnetic field and that the parallel component of the group velocity is closely approximated by the group velocity for longitudinal propagation. It is found that for frequencies near the proton gyrofrequency</p>			

14. KEY WORDS	LINK A		LINK B		LINK C	
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